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Title: Biogeochemical applications of a newly developed carbocation force field: Published in: None (EN). ISBN Author: Van Duin, A.C.T. Thesis advisor.

As the fluid moved, the magnetic field would go with it. The theorem describing this effect is called the frozen-in-field theorem. Even in a fluid with a finite conductivity, new field is generated by stretching field lines as the fluid moves in ways that deform it. This process could go on generating new field indefinitely, were it not that as the magnetic field increases in strength, it resists fluid motion. The temperature increases towards the center of the Earth, and the higher temperature of the fluid lower down makes it buoyant. This buoyancy is enhanced by chemical separation: As the core cools, some of the molten iron solidifies and is plated to the inner core. In the process, lighter elements are left behind in the fluid, making it lighter. This is called compositional convection. A Coriolis effect, caused by the overall planetary rotation, tends to organize the flow into rolls aligned along the north-south polar axis. Early in its history the Sun went through a T-Tauri phase in which the solar wind would have had a magnetic field orders of magnitude larger than the present solar wind. An alternative source is currents in the core-mantle boundary driven by chemical reactions or variations in thermal or electric conductivity. Such effects may still provide a small bias that are part of the boundary conditions for the geodynamo. Simulation of the MHD equations is performed on a 3D grid of points and the fineness of the grid, which in part determines the realism of the solutions, is limited mainly by computer power. For decades, theorists were confined to creating kinematic dynamo computer models in which the fluid motion is chosen in advance and the effect on the magnetic field calculated. Kinematic dynamo theory was mainly a matter of trying different flow geometries and testing whether such geometries could sustain a dynamo. Such a field is always generated near where the atmosphere is closest to the Sun, causing daily alterations that can deflect surface magnetic fields by as much as one degree. Typical daily variations of field strength are about 25 nanoteslas nT one part in 10^5 , with variations over a few seconds of typically around 1 nT one part in 50, Such observatories can measure and forecast magnetic conditions such as magnetic storms that sometimes affect communications, electric power, and other human activities. The military determines local geomagnetic field characteristics, in order to detect anomalies in the natural background that might be caused by a significant metallic object such as a submerged submarine. Commercially, geophysical prospecting companies also use magnetic detectors to identify naturally occurring anomalies from ore bodies, such as the Kursk Magnetic Anomaly. Using magnetic instruments adapted from airborne magnetic anomaly detectors developed during World War II to detect submarines, [64] the magnetic variations across the ocean floor have been mapped. Basalt "the iron-rich, volcanic rock making up the ocean floor [65]" contains a strongly magnetic mineral magnetite and can locally distort compass readings. The distortion was recognized by Icelandic mariners as early as the late 18th century. If an accurate estimate of the field at some other place and time is needed, the measurements must be converted to a model and the model used to make predictions. Multipole expansion Schematic representation of spherical harmonics on a sphere and their nodal lines. Example of a quadrupole field. This can also be constructed by moving two dipoles together. This was first done by Carl Friedrich Gauss. They are the product of two functions, one that depends on latitude and one on longitude. The function of longitude is zero along zero or more great circles passing through the North and South Poles; the number of such nodal lines is the absolute value of the order m . Each harmonic is equivalent to a particular arrangement of magnetic charges at the center of the Earth. A monopole is an isolated magnetic charge, which has never been observed. A dipole is equivalent to two opposing charges brought close together and a quadrupole to two dipoles brought together. A quadrupole field is shown in the lower figure on the right. A magnetic field is a vector field, but if it is expressed in Cartesian components X, Y, Z , each component is the derivative of the same scalar function called the magnetic potential. The next three coefficients " $g_{10}, g_{11},$ and h_{11} " determine the direction and magnitude of the

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dipole contribution. The increasing terms fit the external sources currents in the ionosphere and magnetosphere. However, averaged over a few years the external contributions average to zero. The radius of the outer core is about half of the radius of the Earth. If the field at the core-mantle boundary is fit to spherical harmonics, the dipole part is smaller by a factor of about 8 at the surface, the quadrupole part by a factor of 16, and so on. Thus, only the components with large wavelengths can be noticeable at the surface. From a variety of arguments, it is usually assumed that only terms up to degree 14 or less have their origin in the core. Smaller features are attributed to crustal anomalies. It is updated every five years. Subsequent models are truncated at degree 13 coefficients. This model truncates at degree 12 coefficients with an approximate spatial resolution of 3, kilometers. It was compiled from satellite, marine, aeromagnetic and ground magnetic surveys. These include AM radio signals and ordinary electronic equipment that might be found in businesses or private homes.

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